

Surface elevation changes on Lachman Crags ice caps (north-eastern Antarctic Peninsula) since 1979 indicated by DEMs and ICESat data

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ABSTRACT. The long-term warming on the Antarctic Peninsula in the second half of the 20th century prompted rapid retreat of glaciers on the peninsula and surrounding islands. Retreat accelerated until the beginning of the new millennium when the regional warming trend significantly decreased. The response of glaciers to the change in temperature trend has been observed around the northern part of the Antarctic Peninsula but the timing of the shift from the surface lowering to mass gain remains unclear. Using historical aerial photographs, DEMs and satellite altimeter data from ICESat, we estimate areal and surface elevation changes of two small ice caps in the northern part of James Ross Island over the last 39 years. The glacierized area on Lachman Crags decreased from 4.337 ± 0.037 to 3.581 ± 0.014 km² (−17.4%) between 1979 and 2006 and then increased to 3.597 ± 0.047 km² (0.4%) until 2016. Surface lowering observed on ice caps after 1979 continued at least until 2008 as indicated by the ICESat data. The change from the lowering trend to increase in glacier surface elevation probably occurred after the ablation season 2008/09, which ranks among the warmest summers in the north-eastern Antarctic Peninsula since the mid-20th century.

KEYWORDS: Antarctic glaciology, climate change, ice cap, ice thickness measurements, laser altimetry

1. INTRODUCTION

The Antarctic Peninsula (AP) and the sub-Antarctic islands around its northern tip represent a large glacierized area with numerous glaciers and ice caps outside the Antarctic ice sheet. Most of these glaciers are considered to be in retreat and losing mass, but available glaciological observations are sparse and regionally biased. Changes in length, area, volume or mass were determined only for fraction of the local glaciers, and the AP is ranked among poorly sampled regions (Mernild and others, 2013; Zemp and others, 2015). Most regional studies of glacier volume/mass changes have used topographic elevation data derived from stereoscopic aerial photographs (Rivera and others, 2005b; Molina and others, 2007; Engel and others, 2012; Kunz and others, 2012; Osmanoglu and others, 2014; Pełlicki and others, 2017) and satellite imagery (Rückamp and others, 2011; Davies and others, 2012; da Rosa and others, 2015; Fieber and others, 2018), both of which have limited vertical resolution and time coverage. Since the beginning of 21st century, in situ glaciological monitoring programmes have been initiated to determine the mass balance of small ice caps and glaciers around the AP. In summer 1999–2000, mass-balance measurements started on Bahía del Diablo Glacier on Vega Island on the eastern side of the AP (Fig. 1), with data having been reported to the World Glacier Monitoring Service (WGMS). Additional monitoring programmes were initiated on Livingston Island (Navarro and others, 2013) and King George Island (e.g., Sobota and others, 2015; Mavlyudov, 2016), South Shetland Islands (SSI), ~130 km north-west of the northern tip of the AP. Recently, a surface mass-balance (SMB)

record has been reported for small glaciers on James Ross Island (JRI) on the eastern side of the AP (Engel and others, 2018). Nevertheless, the WGMS database contains mass-balance data only for three glaciers from the AP region.

The AP belongs to the regions that experienced the largest atmospheric warming on Earth over the second half of 20th century. The mean annual air temperature increased by ~0.5°C/decade around the AP (Turner and others, 2005), causing the collapse of ice shelves along its western and north-eastern coast (Cook and Vaughan, 2010) and accelerating the retreat of glaciers (Davies and others, 2012; Scambos and others, 2014) and ice sheets (Wouters and others, 2015). The temperature increase decelerated at the beginning of the new millennium as indicated by a significant decrease in the warming trends reported from the AP region (e.g., Turner and others, 2016). Recent analyses of temperature records from JRI (Ambrožová and Láška, 2016) and the northern AP (Oliva and others, 2017) have even shown that the regional warming trend turned to a cooling during 1999–2015. The response of glaciers to the change in the temperature trend has been observed in the northern part of the AP (Oliva and others, 2017 and references therein) but the timing of the shift from predominantly negative annual SMB to mass gains remains unclear. The first indication of the change in the mass-balance trends over the northern AP was reported by Navarro and others (2013) for Hurd and Johnsons glaciers on Livingston Island (SSI) for the mass-balance year 2008. A later onset of the net mass gain was observed subsequently in the glaciological record for Bellingshausen Ice Dome on King George Island (SSI), which experienced a positive annual SMB in 2010

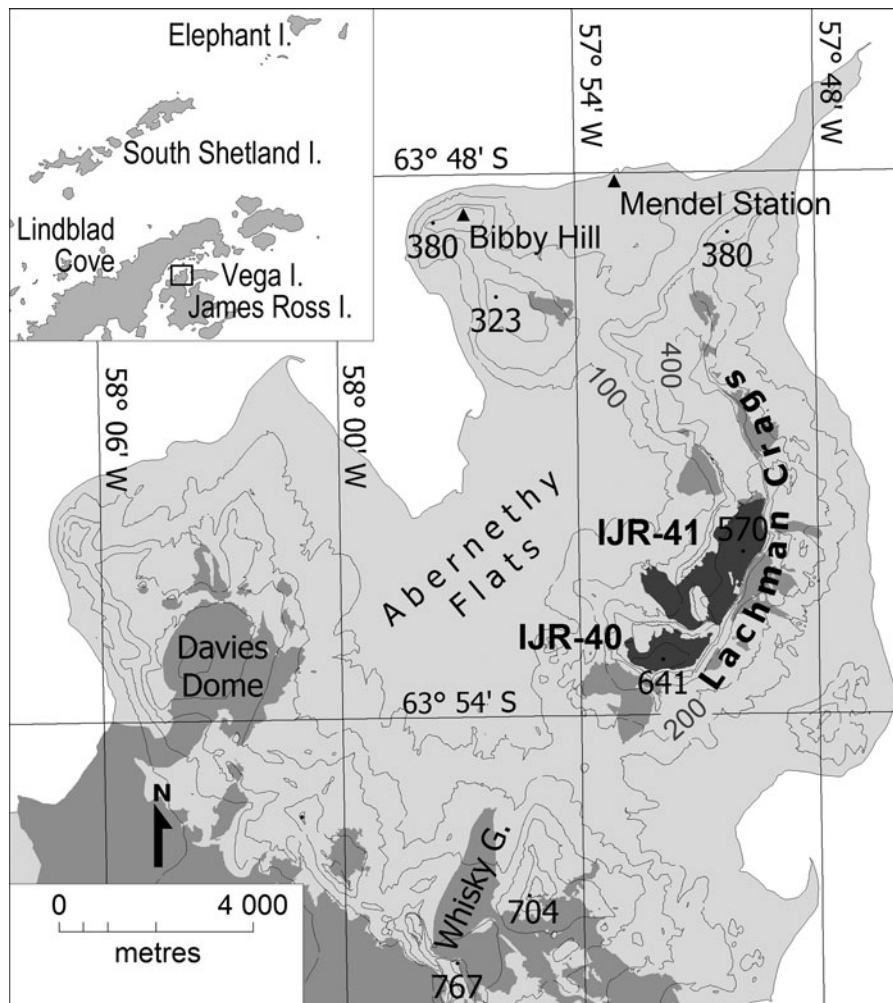


Fig. 1. Location of glaciers (medium grey) and investigated ice caps (dark grey) on Lachman Crags in the northern part of James Ross Island. The inset shows the location of the study area in the north-eastern part of the Antarctic Peninsula.

(Mavlyudov, 2014). The later timing of the shift was proved by the mass-balance dataset for Bahía del Diablo Glacier on Vega Island (Marinsek and Ermolin, 2015), which represents the north-eastern part of the AP. The recently published data from Whisky or Lookalike Glacier (cf. CGS, 2009; SCAR, 2019) and Davies Dome on JRI (Fig. 1) have confirmed predominantly positive annual SMB since the mass-balance year 2010 (Engel and others, 2018) but the time coverage of the data precludes the determination of a possible earlier shift in the mass-balance trend.

In order to determine the possible earlier initiation of the annual mass gain on JRI, we analyse the 1979 and 2006 digital elevation models (DEMs) and the Ice Cloud and Land Elevation Satellite (ICESat) data. The satellite altimetry data provide the information on the surface elevation changes of the IJR-41 ice cap on Lachman Crags (LC) plateau in the northern part of JRI (Fig. 1) over the period 2003–2008. The DEMs based on 1979 and 2006 aerial photographs collected by the British Antarctic Survey (BAS) allow us to analyse the surface elevation changes on this ice cap back to 1979. Moreover, we use the ICESat data, DEMs and static Differential Global Navigation Satellite System (DGNSS) measurement for ICESat ground tracks to access the recent trends in the surface elevation changes on the ice cap. Finally, we determine glacier ice thickness along ICESat tracks using the ground penetrating radar

(GPR) and we estimate the time of complete ice decay based on the ice thickness data and observed the rate of glacier surface lowering.

2. STUDY AREA

JRI is situated on the eastern flank of the northern part of the AP. The island extends between 63°46' and 64°27'S, and reaches 1600 m at the summit of Mount Haddington. The glacial cover of JRI consists of the Mount Haddington ice cap, minor ice caps and outlets at its periphery and small individual glaciers on the Ulu Peninsula in the northern part of the island. The surface of the peninsula is distinguished by cliff-bounded plateaux landforms rising up to ~600 m a.s.l. The prominent flat-topped forms (mesas) are controlled by the bedrock geology, with the late Neogene basalt lava flows forming resistant caps on a Cretaceous sedimentary strata (Hambrey and others, 2008). These mesas are interpreted as the remnants of voluminous lava-fed deltas sourced from a vent at Mount Haddington stratovolcano (Smellie and others, 2008). The LC consists of the thickest lava caprock of any delta on JRI (Smellie and others, 2008), and it dominates the landscape of the north-eastern part of the Ulu Peninsula (Figs 1, 2).

The western side of LC borders low-angle slopes to the Abernethy Flats rising from ~30 to 380–600 m a.s.l.

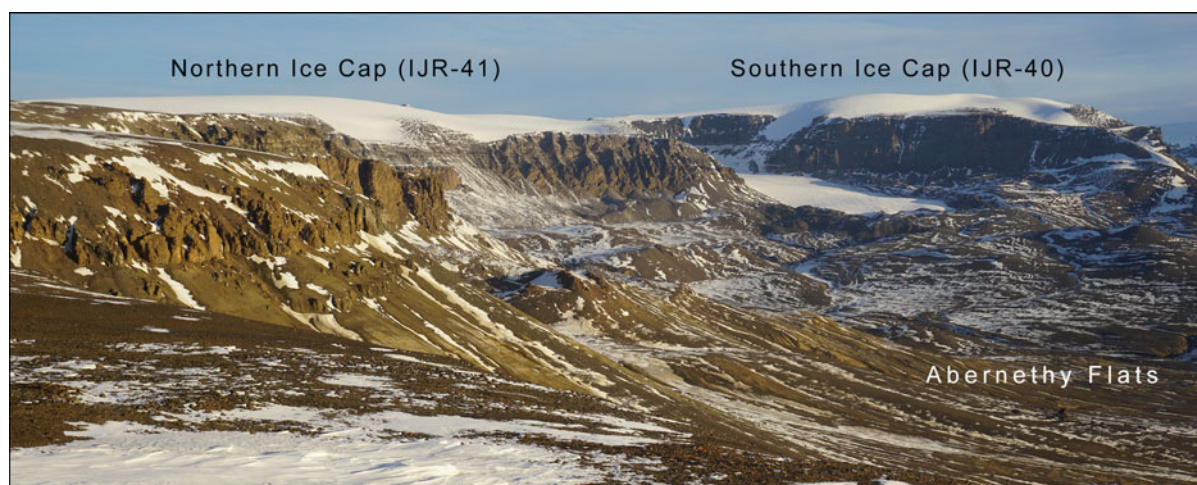


Fig. 2. View of the western side of Lachman Crags from the north-west. Photograph taken on 10 March 2018 from the southern tip of Johnson Mesa.

(Fig. 2). This side is mantled with extensive scree and numerous protalus ramparts that are missing only in sections where two small glaciers descend from the steep cliffs of LC. Scree slopes and protalus ramparts are also formed along the north-western face of LC, east of two abandoned cirques with a steep headwall and overdeepened basins. The eastern and southern face of LC is distinctly different from the western slopes. The straight margin of the LC is sharply delimited by vertical cliffs, which increase in height towards the south from ~ 130 to 300 m. Perennial snow banks and protalus ramparts lie at the foot of a lower northern section of the face, while large mass movement blocks, glaciers and rock glaciers dominate along the higher part of LC. The mass movement blocks separated from the mesa form chaotic hummocky terrain that descends in steps to the seashore.

The surface of LC is inclined to the north, descending from 600 m a.s.l. at its southern tip to 380 m a.s.l. at the northern end. The northern part of the mesa (380–430 m a.s.l.) is ice-free and its flat surface is covered with blockfields and sorted polygon nets. Seasonal snow cover is effectively transported from its surface by the prevailing south-westerly winds (Hrbáček and others, 2016), feeding perennial snow fields and glaciers in the lee-side of the mesa. Two ice caps cover the southern part of LC (Fig. 2). A larger ice cap (referred as IJR-41 in the inventory by Rabassa and others, 1982) is elongated in the NNE direction being 2.8 km long ($63^{\circ}51'34''$ – $63^{\circ}53'2''$ S, $57^{\circ}49'17''$ – $57^{\circ}52'29''$ W) and having an area of 2.7 km². A prominent south-west facing outlet drains the central part of the IJR-41 ice cap with a flat top at 571 m a.s.l. This outlet joins a small glacier snout, which forms a large cirque between the IJR-41 and IJR-40 ice caps. The snout flows towards north-west and terminates at the elevation of 196 m a.s.l. A smaller IJR-40 ice cap with an area of 0.9 km² covers the southernmost part of LC ($63^{\circ}53'2''$ – $63^{\circ}53'31''$ S, $57^{\circ}50'40''$ – $57^{\circ}52'50''$ W). It is located at the elevation of 470–649 m a.s.l. and its 2.3 km long axis has the ENE direction. Both ice caps represent an accumulation area of small stagnant glaciers, which are located below the eastern face of the mesa. These glaciers are detached from the ice caps and thus gain mass due to the snowdrift from the ice caps by winds. San José and Lachman (IJR-38) glaciers at the foot of the southern face of LC and Triangular glacier in its western slope are less influenced by the wind-driven snow redistribution.

The climate of JRI is influenced by the 1000–1800 m high orographic barrier of the AP (King and others, 2003) and the Southern Annular Mode (Marshall and others, 2006). A positive trend in this annular mode increases north-westerly air flow over the AP and amplifies warming on the eastern side of the AP due to the föhn effect (Grosvenor and others, 2014). The mean annual air temperatures (2005–15) in the northern part of JRI range from -7.0°C at the sea level (Mendel Base, 10 m a.s.l.) to lower than -8.0°C (Bibby Hill, 375 m a.s.l.) in higher-elevated regions (Ambrožová and Láška, 2016). Positive air temperatures over the short summer season (2–3 months) cause snow and ice to melt on glaciers, these effects vary strongly in response to cloudiness and solar radiation. The modelled annual precipitation is estimated between 200 and 500 mm w.e. (van Lipzig and others, 2004). Since the mid-20th century, snowfall and summer melt intensity has significantly increased (Abram and others, 2013; Medley and Thomas, 2019).

3. DATA AND METHODS

The BAS aerial photographs taken in 1979 and 2006 and WorldView-2 images from 2016 (Table 1) were used to extract information for an extent of investigated glaciers. The aerial photograph-based DEMs produced by the GEODIS company (the Czech Republic) were used to estimate ice elevations and thinning rates over the glacier surface between 1979 and 2006. The dataset has rms errors of 2.0 m in both the horizontal and vertical directions (1979) and 0.7 m horizontally and 0.8 m vertically (2006), respectively (Meixner, 2009). The accuracy of DEMs derived by stereo-processing over glaciers often suffers from a lack of surface patterns resulting in poor quality elevation information (Rivera and others, 2005a). We checked the original aerial photographs and found sufficient surface-patterned textures on the glacier surface. We thus assume that the DEM quality over glaciers is not lower than over the rest of the dataset. Surface elevation changes and annual thinning rates after 2006 were calculated using the 2006 DEM and the Reference Elevation Model of Antarctica (REMA) created by the Polar Geospatial Center (PGC, 2018). REMA was constructed from high-resolution (~ 0.5 m) stereoscopic imagery collected by DigitalGlobe, Inc. constellation of optical imaging satellites (PGC, 2018).

Table 1. Summary of remote-sensing data

Acquisition date	Camera/spacecraft	Images/bands	Pixel size at ground m	Image identification
January 1979	Vinten 70 mm	panchromatic	2.1	BAS/JR/1/79
February 14 2006	Leica RC30	panchromatic	0.52	BAS/RN/06
December 3 2016	WorldView-2	Pan-MS1-MS2	0.61	10300100626DD300

This model portrays first-return elevation values that reflect temporally variable conditions such as snow cover. Snow depth records and field observations on the Ulu Peninsula indicate irregular distribution of snow cover, with a maximum depth of fresh snow ranging from 0.3 m at the sea level (Hrbáček and others, 2016) to ~ 0.6 m on glaciers. The effect of variable snow depth on the accuracy of the DEMs cannot be excluded but it remains on the order of magnitude lower than vertical uncertainty of DEM-derived elevation values. Glacier-elevation changes on the flat top surface of IJR-41 ice cap were estimated in more detail using the available DEMs, ICESat data acquired during the period 2003–2008 and in situ DGNSS measurements from 2018.

The ground track of ICESat satellite crosses LC providing elevation measurements of land surface above the TOPEX-Poseidon ellipsoid along the satellite ground track each 172 m. Each point measurement corresponds to a footprint with 70 m in diameter. Due to energy management constraints, the laser was on during limited periods only, mainly in March, May and October. The ground tracks are fixed but single flyovers do not follow exactly the same line on the surface. In order to derive the change in ice surface elevation for small glaciers, ICESat elevations are related to a static elevation reference such as SRTM DEM (e.g., Gardner and others, 2013; Neckel and others, 2014). We used the static DGNSS measurements at the ICESat footprints as a reference that increases the accuracy of the determined elevation compared with the DEM reference (Kropáček and others, 2014). We used the geodetic records of GLA14 product release 34. The surface elevation was determined at 72 individual locations spread along six ICESat ground tracks over ~ 2 km on the ice dome. Erroneous measurements due to clouds were identified by the elevation threshold of 100 m and removed.

DGNSS technology was applied to determine the contemporary ice surface elevation for the ICESat ground tracks. The data for 72 points were collected during a field campaign in January 2018 using a Trimble Geo7X roving station (Trimble Navigation Limited, 2013). The reading lasted 3 min at each site to obtain centimetre accuracy after post-processing. The coordinate points were corrected in relation to the base station mounted on a stable surface near the J.G. Mendel Czech Antarctic Station ($63^{\circ}48'S$; $57^{\circ}53'W$) and the correction data acquired by the German Antarctic Receiving Station at the Chilean station General Bernardo O'Higgins ($63^{\circ}19'S$; $57^{\circ}54'W$). The data collected on LC ice cap were post-processed for differential correction using the Trimble GPS Pathfinder Office software.

GPR was used to determine the ice thickness for ground tracks (55 points) of ICESat on IJR-41 ice cap. The GPR data were collected in February 2017 along five parallel profiles on the dome (Fig. 3). GPR profiling was carried out using an unshielded 50 MHz Rough Terrain Antenna and RAMAC CU-II control unit (MALÁ GeoScience, 2005). The scan spacing was set to 0.1 s and the GPR was man-hauled at a

mean speed of ~ 3 km h^{-1} , yielding one radar trace every ~ 0.01 m. A Garmin GPSMAP 60CSx receiver (Garmin International, 2007) was used to determine the coordinates for the corresponding trace numbers with the horizontal position accuracy < 5 m. The GPR data were processed and interpreted using the ReflexW software version 8.5 (Sandmeier, 2017). Two-way travel times to depths were converted assuming a radio-wave velocity of 168 m μs^{-1} reported by Narod and Clarke (1994) for cold glacier ice. The raw GPR profiles were filtered using an automatic gain function (AGC-Gain), background removal and static (topographic) correction. The glacier bed was selected on radargrams by a phase follower tool in ReflexW. The automatic assignment of picks was controlled and eventually redirected to the desired phase at the glacier bed. The picks were collected only for the bed reflections that can be clearly identified in the radargrams.

The uncertainty of ice thickness determined for ICESat points was estimated following the approach proposed by Lapazaran and others (2016), who describes the error in thickness (i) retrieved from the GPR measurement and (ii) due to the incorrect horizontal positioning of the measurement point. For the calculations of the first uncertainty component, we considered a conservative estimate of vertical resolution (a half of the signal's wavelength in ice, i.e.

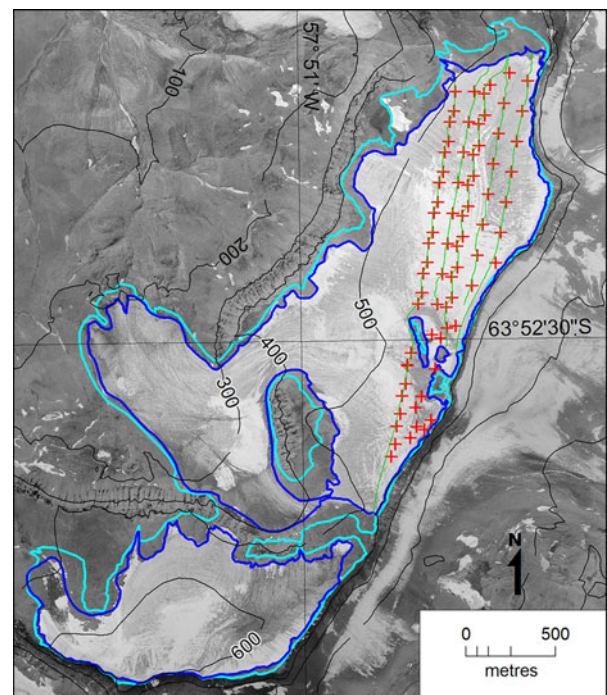


Fig. 3. Ground tracks of ICESat (red crosses) and GPR profiles (green lines) on the flat-topped surface of IJR-41 ice cap superimposed on an aerial photography acquired in February 2006. Light and dark blue lines represent the glacierized area in 1979 and 2016, respectively.

1.68 m for the 50 MHz antenna) and a relative error in radio-wave velocity of 2.4%, which corresponds to a range of velocity values from 166 to 170 m μs^{-1} . We chose this relatively narrow range since we anticipated only minor variations in ice and snow characteristics (density, water/air content) on the investigated part of LC ice cap (500–570 m a.s.l.). The positioning-related ice-thickness error was determined from the variability of ice thickness along and across the profiling direction. Taking into account that the radar traces were positioned using uncorrected GPS data, we evaluated the maximum absolute value of the differences between the value at the ICESat point and the neighbouring values within a circle with a radius of 5 m. The variability of ice thickness along and across the profiling direction was calculated from the values picked for radar traces and interpolated between GPR profiles, respectively.

4. RESULTS

The area and ice-surface elevation changes determined for the investigated glaciers are given in Table 2. A single ice cap with the area of $4.337 \pm 0.037 \text{ km}^2$ covered LC in 1979. Its surface decreased significantly before 2006, losing 17.4% of its area. Assuming a constant rate of change, the ice cap therefore lost $0.029 \text{ km}^2 \text{ a}^{-1}$ during the period 1979–2006. Most of the measured areal loss occurred at the southern part of LC at 450–600 m a.s.l. (Fig. 3). As a result, ice cover on LC diverged on two separate ice caps with the area of 2.709 ± 0.009 and $0.872 \pm 0.004 \text{ km}^2$ on the central and southern parts of LC, respectively. From 2006 to 2016, the surface area of the larger of the two ice caps (IJR-41) remained almost the same. The glacier decreased to $2.674 \pm 0.031 \text{ km}^2$, losing only 1.3% of its area at a mean rate of $0.003 \text{ km}^2 \text{ a}^{-1}$. By contrast, the surface area of IJR-40 increased to $0.922 \pm 0.015 \text{ km}^2$ over the same period, gaining 5.8% of its area at a mean rate of $0.005 \text{ km}^2 \text{ a}^{-1}$.

The glacierized area on LC has experienced only minor surface elevation changes since 1979. The surface lowering observed along western margin of the glaciers was partly compensated by an increase in surface elevation above 560 m. Between 1979 and 2006, the mean surface elevation of IJR-41 ice cap decreased by $7.3 \pm 2.2 \text{ m}$, at a mean rate of -0.28 m a^{-1} (Table 2). By contrast, the mean surface elevation of IJR-40 ice cap increased by $5.6 \pm 2.2 \text{ m}$, with a mean rate of 0.21 m a^{-1} . The discrepancy in surface

elevation changes may be attributed to the recession of the lower-elevated periphery of IJR-40 ice cap and more prominent accumulation on its upper part that is situated $\sim 70 \text{ m}$ higher than the top of IJR-41. Both ice caps experienced an increase in surface elevation over the period 2006–2015. Mean surface elevations of IJR-41 and IJR-40 ice caps increased on average by 3.9 ± 0.8 and $1.2 \pm 0.8 \text{ m}$, respectively. On average, mean surface elevations increased at a rate of 0.43 and 0.13 m a^{-1} .

The elevation data determined at 72 individual ICESat points document surface elevation changes on the flat top part of IJR-41 ice cap over the period 1979–2018 (Table 3). The mean surface elevation of the dome decreased by 4.0 m between 1979 and 2006 at a mean lowering rate of -0.1 m a^{-1} . Only the highest part of the dome (26 out of 72 ICESat points) experienced an increase in surface elevation by 3.7 m on average over this period. From 2006 to 2015, the surface elevation of the dome increased on average by 4.0 m (a mean rate of 0.4 m a^{-1}) with an increase observed at 63 points. Since 2015, the dome surface has returned to the lowering trend, which is recorded at all points. From 2015 to 2018, the surface elevation of the dome decreased by 3.7 m at a mean rate of -1.2 m a^{-1} . The surface lowering prevailed on the dome over most of the 1979–2018 as indicated by a net mean decrease value of 3.7 m and a mean lowering rate of -0.1 m a^{-1} . The decreasing trend was also confirmed for the period 2003–2008 based on the ICESat data (Table 4 and Fig. 4). An overall increase in surface elevation was restricted only to the highest part of the dome (Fig. 5).

The mean glacier ice thickness for 55 points on the flat top surface of IJR-41 ice cap is $43.1 \pm 1.7 \text{ m}$ (Table 3). The greatest thickness was observed on the highest part of the dome, where the glacier ice was as much as $71.4 \pm 2.1 \text{ m}$ thick. The ice thickness decreases slowly towards the northern glacier margin, which terminates on the slightly inclined bedrock at the central part of the LC mesa (Figs 6, 7). The thickness decreases more abruptly in the southern part of the dome where steep bedrock outcrops protrude above ice surface (Fig. 6). In the west-east direction, greater ice thickness values (up to 60 m) were recorded along the western ICESat ground tracks. This part of the dome represents a source area for two lobes that descend towards the western margin of the ice cap. The ice thickness ranges from 10 to 40 m along the eastern margin of the dome, which forms an ice-cliff above the basalt and hyaloclastite

Table 2. Changes of surface area and elevation of ice caps on Lachman Crags since January 1979

Year	Surface area km^2		Mean elevation m a.s.l.	
	IJR-41	IJR-40	IJR-41	IJR-40
1979		4.337 ± 0.037		490.3 ± 10.4
	3.219 ± 0.027	1.117 ± 0.012	457.4 ± 10.3	585.4 ± 6.1
2006	2.709 ± 0.009	0.872 ± 0.004	450.1 ± 10.7	591.0 ± 5.6
2016*	2.674 ± 0.031	0.922 ± 0.015	454.0 ± 10.7	592.2 ± 5.7
1979–2006				
Change	-0.511 ± 0.028 (–15.9%)	-0.245 ± 0.013 (–21.9%)	-7.3 ± 2.2 (–1.6%)	5.6 ± 2.2 (+1.0%)
Mean annual change	-0.020 ± 0.001 (–0.6%)	-0.0094 ± 0.0005 (–0.8%)	-0.28 ± 0.08 (–0.06%)	0.21 ± 0.08 (+0.04%)
2006–2016				
Change	-0.035 ± 0.033 (–1.3%)	0.050 ± 0.016 (+5.8%)	3.9 ± 0.8 (+0.9%)	1.2 ± 0.8 (+0.2%)
Mean annual change	-0.003 ± 0.003 (–0.1%)	0.005 ± 0.001 (+0.5%)	0.43 ± 0.09 (+0.1%)	0.13 ± 0.09 (+0.02%)

* Mean elevation determined for 2015 (PGC, 2018).

Table 3. Summary of elevation data and ice thickness at ICESat points. A dash indicates unknown ice thickness at points that were not covered with GPR survey. The decay time of ice is determined from the surface-lowering trend observed during the period 1979–2018 (Trend 1) and 2015–2018 (Trend 2)

Date of ICESat tracks	Latitude (°S)	Longitude (°W)	Elevation m a.s.l.					Ice thickness m	Elevation change m a ⁻¹		Ice decay time y	
			2003–2008	1979	2006	2015	2018		1979–2018	2015–2018	Trend 1	Trend 2
6-11-2003	63.862376	57.832055	498.4	509.6	511.8	501.6	496.9	7.4 ± 1.9	-0.33	-1.55	22	5
6-11-2003	63.863897	57.832702	521.8	532.5	517.8	522.8	517.4	19.6 ± 1.6	-0.40	-1.80	50	11
6-11-2003	63.865419	57.833352	534.8	548.2	532.0	538.0	534.0	37.0 ± 1.5	-0.37	-1.33	99	28
6-11-2003	63.866942	57.834003	543.4	555.1	542.7	549.1	545.3	47.8 ± 1.7	-0.26	-1.25	185	38
6-11-2003	63.868465	57.834655	556.1	560.0	552.7	560.3	556.5	59.5 ± 1.5	-0.09	-1.25	644	47
6-11-2003	63.869989	57.835308	564.3	561.6	558.7	568.0	564.2	61.8 ± 2.1	0.07	-1.28	-	48
6-11-2003	63.871514	57.835963	564.9	560.6	560.8	568.6	564.4	58.5 ± 1.9	0.10	-1.40	-	42
6-11-2003	63.873039	57.836619	553.1	554.7	548.6	556.5	552.2	31.4 ± 1.9	-0.06	-1.41	483	22
6-11-2003	63.876090	57.837932	522.2	523.0	523.6	527.9	523.6	-	0.02	-1.42	-	19
6-11-2003	63.877611	57.838587	531.6	538.4	532.4	536.1	532.8	-	-0.15	-1.09	267	36
6-11-2003	63.879131	57.839240	542.3	550.0	540.0	545.3	541.7	-	-0.22	-1.18	217	40
6-11-2003	63.880651	57.839895	539.3	545.9	535.6	541.2	537.1	-	-0.23	-1.38	199	33
8-6-2004	63.862075	57.828092	524.3	529.5	533.9	523.6	519.1	28.5 ± 1.1	-0.28	-1.52	103	19
8-6-2004	63.863592	57.828757	542.3	548.0	541.1	540.9	536.4	40.6 ± 1.7	-0.30	-1.51	133	27
8-6-2004	63.865110	57.829424	551.4	557.2	546.7	550.3	546.3	34.7 ± 1.1	-0.29	-1.33	121	26
8-6-2004	63.866632	57.830093	557.8	560.0	555.2	557.9	554.6	56.8 ± 1.8	-0.14	-1.10	398	52
8-6-2004	63.868155	57.830762	565.5	560.0	561.5	567.5	564.1	51.4 ± 1.9	0.11	-1.13	-	46
8-6-2004	63.869680	57.831428	571.7	560.0	565.0	574.3	571.0	64.7 ± 2.4	0.29	-1.09	-	60
8-6-2004	63.871206	57.832089	571.9	559.4	565.0	575.0	572.0	64.1 ± 1.6	0.33	-0.97	-	66
8-6-2004	63.872733	57.832739	562.6	554.3	558.9	566.1	562.7	60.3 ± 2.1	0.22	-1.12	-	54
8-6-2004	63.874261	57.833385	548.7	552.0	545.9	550.0	546.3	22.8 ± 2.9	-0.15	-1.23	153	19
8-6-2004	63.878844	57.835318	539.8	545.5	535.8	538.6	535.6	-	-0.26	-0.99	16	4
11-3-2005	63.878212	57.837057	540.7	549.4	539.2	542.2	539.5	-	-0.26	-0.93	76	21
11-3-2005	63.879741	57.837715	543.5	550.0	540.0	545.0	541.4	-	-0.23	-1.20	82	16
10-6-2005	63.863353	57.832229	516.2	529.8	518.5	517.6	511.9	17.8 ± 1.4	-0.47	-1.90	38	9
10-6-2005	63.864879	57.832884	531.7	545.2	533.5	534.8	530.5	27.8 ± 2.2	-0.39	-1.45	72	19
10-6-2005	63.866401	57.833539	541.6	554.7	541.5	546.5	542.6	43.2 ± 1.7	-0.32	-1.29	136	34
10-6-2005	63.867921	57.834196	552.9	557.7	550.9	557.5	553.7	56.3 ± 1.6	-0.11	-1.25	531	45
10-6-2005	63.869439	57.834852	563.4	560.0	558.7	566.6	562.9	59.8 ± 1.6	0.08	-1.25	-	48
10-6-2005	63.870955	57.835508	566.6	564.2	564.7	570.2	566.4	55.6 ± 1.5	0.06	-1.28	-	44
10-6-2005	63.872470	57.836164	558.9	557.7	555.6	563.7	560.1	51.6 ± 2.2	0.06	-1.23	-	42
10-6-2005	63.875495	57.837475	524.1	530.4	522.6	530.2	525.4	-	-0.13	-1.59	113	9
10-6-2005	63.877006	57.838129	528.1	531.0	526.5	533.1	529.3	-	-0.04	-1.25	895	31
10-6-2005	63.878517	57.838780	540.3	548.5	537.9	542.4	539.3	-	-0.24	-1.04	169	39
10-6-2005	63.880028	57.839430	542.8	548.7	540.0	545.1	541.3	-	-0.20	-1.30	259	39
11-11-2005	63.862504	57.828822	526.1	534.8	535.7	527.3	522.3	28.8 ± 1.0	-0.33	-1.67	88	17
11-11-2005	63.864028	57.829463	542.5	550.2	540.7	542.9	538.5	35.2 ± 1.4	-0.31	-1.44	115	24
11-11-2005	63.865555	57.830109	551.2	558.2	548.9	551.6	547.9	45.6 ± 1.3	-0.27	-1.26	168	36
11-11-2005	63.867085	57.830762	557.8	560.0	555.4	560.2	557.1	63.5 ± 2.1	-0.08	-1.03	839	61
11-11-2005	63.868620	57.831424	566.2	560.0	564.2	569.7	566.4	48.9 ± 1.5	0.17	-1.10	-	45
11-11-2005	63.870155	57.832091	572.0	560.0	565.0	575.1	571.9	66.7 ± 2.2	0.31	-1.04	-	64
11-11-2005	63.871690	57.832759	570.0	558.9	565.0	573.8	570.8	64.2 ± 1.7	0.31	-0.98	-	65
11-11-2005	63.873223	57.833427	559.0	554.5	555.0	562.4	559.4	50.6 ± 2.2	0.13	-1.01	-	50
11-11-2005	63.874751	57.834092	542.0	550.0	541.6	544.6	541.2	-	-0.23	-1.13	28	6
11-11-2005	63.876276	57.834754	538.4	544.6	536.3	543.4	540.2	-	-0.12	-1.07	0	0
11-11-2005	63.879317	57.836069	541.3	547.5	537.5	541.2	537.6	-	-0.26	-1.21	21	5
14-6-2006	63.861890	57.823792	519.1	526.6	529.8	521.8	518.7	9.3 ± 1.3	-0.21	-1.02	45	9
14-6-2006	63.863411	57.824456	533.3	540.8	542.7	535.0	531.0	12.7 ± 1.6	-0.26	-1.31	50	10
14-6-2006	63.864932	57.825120	542.8	549.8	542.1	543.9	539.7	25.9 ± 1.8	-0.27	-1.42	97	18
14-6-2006	63.866452	57.825781	547.9	556.2	548.1	549.4	546.1	25.9 ± 1.3	-0.27	-1.12	98	23
14-6-2006	63.867969	57.826437	555.6	560.0	552.0	557.9	555.1	40.8 ± 1.4	-0.13	-0.94	316	43
14-6-2006	63.869486	57.827091	561.3	560.0	557.5	563.5	560.7	36.5 ± 2.0	0.02	-0.93	-	39
14-6-2006	63.871003	57.827741	560.0	544.4	556.8	562.8	559.8	37.8 ± 2.9	0.41	-0.97	-	39
23-10-2007	63.861468	57.825878	516.3	524.3	522.9	521.4	517.1	10.8 ± 1.2	-0.19	-1.45	57	7
23-10-2007	63.862994	57.826537	535.3	542.6	543.4	537.7	533.5	32.5 ± 1.7	-0.24	-1.40	135	23
23-10-2007	63.864518	57.827191	545.8	554.5	543.4	548.2	543.7	35.4 ± 1.1	-0.28	-1.52	124	23
23-10-2007	63.866041	57.827843	551.8	560.0	550.3	554.9	550.5	38.6 ± 1.6	-0.25	-1.47	155	26
23-10-2007	63.867564	57.828494	559.5	560.0	557.8	562.8	559.4	47.9 ± 1.4	-0.02	-1.14	2828	42
23-10-2007	63.869087	57.829144	566.9	560.0	564.7	570.1	567.3	47.8 ± 1.4	0.19	-0.96	-	50
23-10-2007	63.870613	57.829795	568.7	557.0	564.5	572.1	569.0	58.6 ± 2.2	0.32	-1.04	-	56
23-10-2007	63.872141	57.830447	560.6	551.3	556.9	565.3	562.0	53.4 ± 1.7	0.28	-1.10	-	49
9-3-2008	63.862391	57.829889	512.4	526.2	532.0	518.4	512.7	25.3 ± 0.9	-0.35	-1.89	72	13
9-3-2008	63.863914	57.830552	532.2	544.6	535.5	537.4	534.0	31.7 ± 1.2	-0.28	-1.15	113	27
9-3-2008	63.865437	57.831210	542.7	556.6	542.4	548.0	544.6	39.1 ± 1.1	-0.32	-1.15	124	34

Table 3. (Cont.)

Date of ICESat tracks	Latitude (°S)	Longitude (°W)	Elevation m a.s.l.					Ice thickness m	Elevation change m a ⁻¹		Ice decay time y	
			2003–2008	1979	2006	2015	2018		1979–2018	2015–2018	Trend 1	Trend 2
9-3-2008	63.866962	57.831862	550.5	560.0	551.9	557.6	553.9	56.0 ± 2.1	-0.16	-1.21	350	46
9-3-2008	63.868487	57.832511	561.0	560.0	561.5	567.4	563.8	55.6 ± 1.4	0.10	-1.20	-	47
9-3-2008	63.870013	57.833156	568.3	560.0	565.0	574.2	570.8	71.4 ± 2.1	0.28	-1.16	-	62
9-3-2008	63.871538	57.833798	568.9	561.0	565.0	574.3	571.1	66.3 ± 1.6	0.27	-1.08	-	61
9-3-2008	63.873064	57.834438	557.9	555.0	557.0	563.5	560.0	48.6 ± 2.7	0.13	-1.15	-	42
9-3-2008	63.874590	57.835075	543.6	552.5	544.7	548.4	544.9	-	-0.20	-1.16	24	4
9-3-2008	63.877643	57.836348	536.5	545.6	538.1	542.6	539.8	-	-0.15	-0.95	51	8
9-3-2008	63.879171	57.836985	541.1	550.0	540.0	544.0	540.9	-	-0.24	-1.05	70	16
		Mean	546.7	549.9	545.9	549.9	546.1	43.10	-0.10	-1.24	233	32
		Std	16.4	11.5	13.3	16.6	17.0	16.45	0.22	0.22	422	18

eastern face of LC. The uncertainties in the ice-thickness values related to the GPR measurement and trace-positioning range from 0.22 to 1.70 m and from 0.42 to 2.84 m, respectively. The total uncertainties range between 0.93 and 2.89 m, with a mean value of 1.71 m and a std dev. of 0.44 m.

5. DISCUSSION

The decrease in the surface area of the ice caps on LC over the 1979–2006 period is within the range of values reported by Engel and others (2012) for small glaciers in the northern part of JRI. The areal loss of 21.9 and 15.9% observed on IJR-40 and IJR-41 ice caps, respectively, coincides with the 22.1% decrease of Davies Dome surface area (Engel and others, 2012). These values are higher than the 10.6% loss reported by Engel and others (2012) for Whisky Glacier, which suggests a more rapid retreat of ice caps than cold-based valley glaciers in this region. This implies higher sensitivity of small ice caps to regional climate settings and local meteorological conditions (e.g., Nesje and others, 2008; Zekollari and others, 2017). The recession of the ice caps in the northern part of JRI is less prominent compared with Subantarctic (*sensu* Cogley and others, 2014) ice caps and their outlets that experience up to an order of magnitude higher retreat and surface-lowering rates (Table 5).

The deceleration of the retreat rate and the increase in surface elevation observed on the investigated ice caps during the 2006–2015 period correspond with the lowered glacier recession around the northern AP (Table 5). Positive

mass-balance values recorded in this region since the glaciological year 2009/10 (Mavlyudov, 2016; Engel and others, 2018; WGMS, 2018) and the low rates of glacier recession in subsequent years (Pełlicki and others, 2017; Pudelko and others, 2018) have been attributed to the deceleration of the regional warming trend and the decrease in summer temperature (Navarro and others, 2013; Lirio and others, 2017; Oliva and others, 2017). The cooling trend in the northern AP has been recently confirmed by Ambrožová and Láška (2016), who found the most prominent decrease in air temperature (1.44°C/decade) over the period 2005–2015 in the northern part of JRI. Moreover, this re-assessment of the temperature records has revealed a significant decrease in summer temperature at AP meteorological stations including Mendel and Bibby Hill stations near LC.

Although glaciers in the AP are influenced by a set of regional and local climatic conditions (e.g., Arigony-Neto and others, 2014), our results confirm the relationship between summer mean air temperatures and surface elevation changes reported by Skvarca and De Angelis (2003) for this region. Surface elevation changes on IJR-41 ice cap indicated by the ICESat data correlate with summer temperatures at Esperanza meteorological station during the period 2003–2008 (Fig. 8). The lowering trend observed on IJR-41 ice cap during the period 2003–2008 (Fig. 4) and surface elevation changes determined from the ICESat tracks measured in June 2006, October 2007 and March 2008 suggest that the glacier lost mass in the mass-balance year 2008. High temperatures recorded in the northern AP during the ablation

Table 4. List of ICESat ground tracks showing the difference of ICESat elevations with respect to the DGNSS measurements from 2018 as median value for each track

Track date	Number of points	Median elevation difference (m)	Standard deviation from mean value	Standard deviation from median value
06 Nov 2003	12	0.50	1.65	1.65
08 Jun 2004	10	2.80	2.17	2.17
11 Mar 2005	2	1.73	0.45	0.45
10 Jun 2005	11	0.22	1.62	1.62
11 Nov 2005	11	0.65	2.01	2.08
14 Jun 2006	7	0.59	1.08	1.27
23 Oct 2007	8	-0.10	1.20	1.27
09 Mar 2008	11	-2.14	1.09	1.11

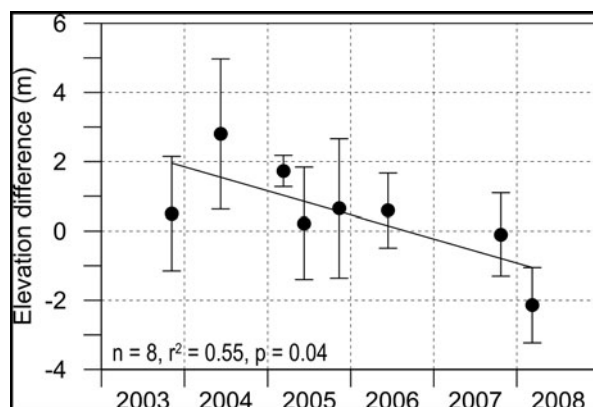


Fig. 4. Median elevation difference of the ICESat tracks with respect to the DGNSS reference.

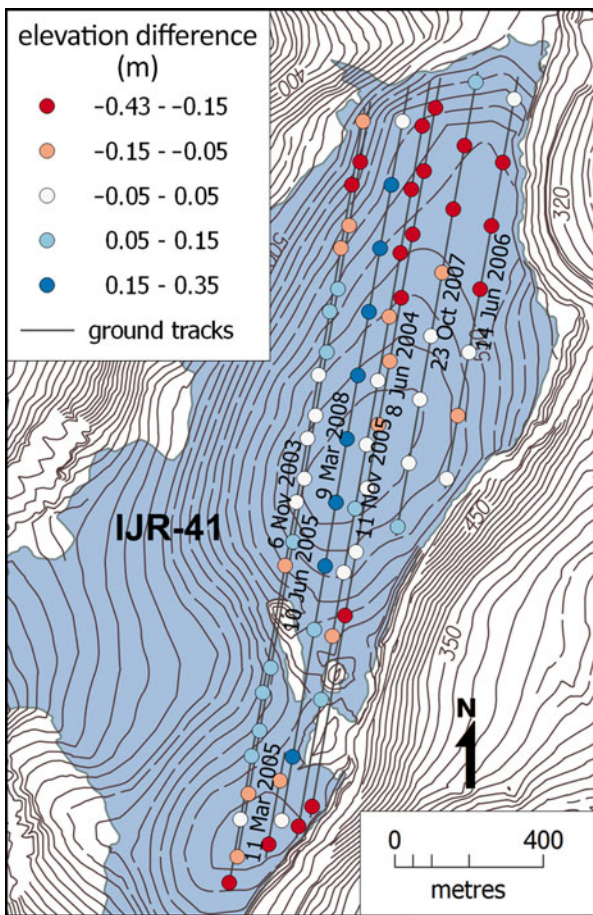


Fig. 5. Median annual difference of ice surface elevation between the ICESat data (2003–2008) and DGNS (2018) measurements.

season 2007/08, which has been the fifth warmest summer (with a mean summer air temperature of 1.7°C) on Esperanza since 1953 according to the Reference Antarctic Data for Environmental Research (READER) database (Turner and others, 2004) supports this view. An increase in surface elevation on the dome in the mass-balance year 2009 cannot be excluded based on the ICESat data as the last available tracks represent the onset of this year. However, warm conditions during the ablation season 2008/09 (1.6°C on Esperanza; READER database) and negative SMB reported for the glaciers around the AP (Fig. 9) make this possibility improbable.

Based on the mean lowering rates and ice thickness determined for 55 points on IJR-41 ice cap, we calculate the time for the complete melting of glacier ice at these points. The melting time must be considered as a rough estimate because the approximation omits future changes in climate conditions and ice dynamics. In order to reflect recent temporal variability of the glacier evolution, we take into account the mean lowering rates for the periods 1979–2018 and 2015–2018. The surface elevation changes over a longer time span range from $+0.41\text{ m a}^{-1}$ at the south-eastern part of the dome to -0.47 m a^{-1} near the north-western ice margin. The calculated ice decay times indicate that the periphery of the ice cap north of the small western outlet will disappear within decades (Fig. 7). The same evolution may be expected near the bedrock outcrops in the southern part of the dome. The melting time increases to centuries towards the central part of the dome that would gain mass under unchanged conditions. The accelerated lowering rates (-0.93 to -1.90 m a^{-1}) observed over the period 2015–2018 reduce the time to complete ice decay roughly by an order of magnitude. If these rates are taken into account,

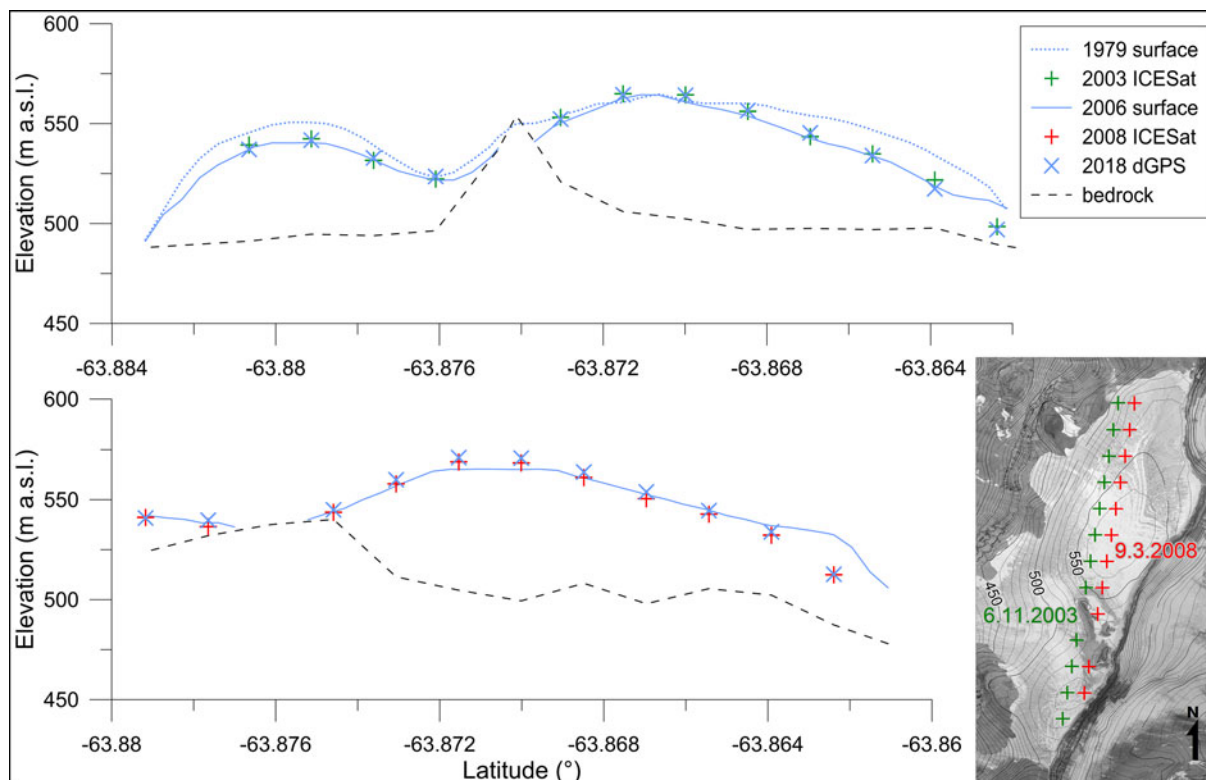


Fig. 6. Glacier-surface elevation changes derived from the DEMs, ICESat data and DGNS measurement on IJR-41 ice cap. The vertical axis is exaggerated by a factor of 4.

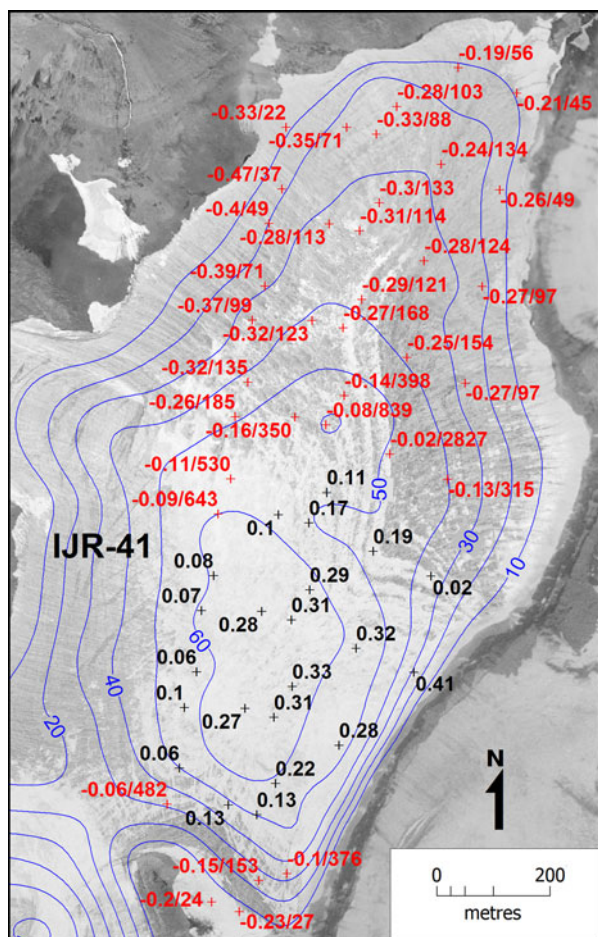


Fig. 7. Ice thickness (blue lines) on the dome and glacier-surface elevation changes (m a^{-1}) at ICESat points (red crosses) over the period 1979–2018. Black figures show the increase in surface elevation while red values indicate surface lowering and years to complete ice melt based on the averaged lowering rates over this period.

even the highest part of the ice cap would disappear in 66 ± 19 years.

The future evolution of ice around the highest part of the dome approximately based on the mean lowering rate for the period 1979–2018 is consistent with the ice-decay estimates reported for glaciers of comparable area in the northern AP region. The estimated melting time of Whisky Glacier and Davies Dome on the Ulu Peninsula is 104 ± 5 and 228 ± 22 years, respectively (Engel and others, 2012). Similarly, Bellingshausen Dome on SSI is predicted to disappear in 285 years if the present warming trend persists (Rückamp and others, 2011). A total decay of IJR-41 ice cap within the next few decades estimated based on the accelerated lowering rate (2015–2018) seems to be unrealistic. A rapid ice-decay during this time span is predicted for small glaciers at low elevation that are located in regions affected by changes in regional atmospheric circulation. Simões and others (2004) described these conditions on the SSI where small glaciers lost up to 83% of their area during the second half of the 20th century. Rapidly receding glaciers that are now close to decay were also reported from South Georgia (Gordon and others, 2008) and the Kerguelen Islands (Berthier and others, 2009).

6. CONCLUSIONS

The analyses of the DEMs, ICESat and DGNSS data revealed a predominant retreat of ice caps on LC since 1979. Between 1979 and 2006, the glacierized area decreased by 17.4% at a mean rate of $0.03 \text{ km}^2 \text{ a}^{-1}$. The observed areal decrease, retreat rate and surface elevation change are in accordance with the values that represent glacier recession in the northern part of JRI, which are by an order of magnitude lower than the values reported for ice caps and outlets on the Subantarctic islands. The glacier retreat on LC decelerated during the period 2006–2016 when the area of IJR-41 ice cap decreased by 1.3% and IJR-40 experienced a 5.8%

Table 5. Mean area and surface elevation changes of ice caps and outlet glaciers along the northern part of the Antarctic Peninsula

Site	Glacier		Period	Area change				Elevation change		Reference
	Name	Type		km^2	%	$\text{km}^2 \text{ a}^{-1}$	%	m	m a^{-1}	
Elephant Island	Doyle	O-m	1957–2014	–	–	–	–	–9.1	–0.16	Fieber and others, 2018
	Helava	O-m	1957–2014	–	–	–	–	–2.1	–0.04	
James Ross Island	Davies Dome	IC-m	1979–2006	–1.840	–22.1	–0.071	–0.8	–8.5	–0.33	Engel and others, 2012 this study
	IJR-40	IC-l	1979–2006	–0.245	–21.9	–0.009	–0.8	+5.6	+0.21	
	IJR-40	IC-l	2006–2016	0.050	5.8	0.005	0.5	+1.2	+0.13	
	IJR-40	IC-l	1979–2006	–0.511	–15.9	–0.020	–0.6	–7.3	–0.28	
King George Island	Whisky	V	1979–2006	–0.284	–10.6	–0.011	–0.4	–10.1	–0.39	Engel and others, 2012 Rückamp and others, 2011
	Bellingshausen Dome	IC-l	1998–2009	–	–	–	–	–	–0.20 to –1.44	
Lindblad Cove	Ecology	O-m	1979–2001	–	–	–	–	–38.4	–1.75	Petlicki and others, 2017 Sobota and others, 2015 da Rosa and others, 2015 Pudelko and others, 2018 Fieber and others, 2018 Fieber and others, 2018
	Ecology/Sphinx	O-m/l	1979–2007	–3.195	–38.0	–0.114	–1.4	–	–	
	Wanda	O-l	1979–2006	–0.690	–30.4	–0.026	–1.1	–	–	
	Warszawa Icefield	IC	1979–2007	–4.810	–24.3	–0.172	–0.9	–	–	
	Warszawa Icefield	IC	1979–2018	–6.100	–30.8	–0.156	–0.8	–	–	
Lindblad Cove	Znosko	O-m	1956–2013	–	–	–	–	–11.1	–0.20	Fieber and others, 2018 Fieber and others, 2018
	Schoeling	O-m	1957–2014	–	–	–	–	+1.7	+0.03	

IC, ice cap; O, outlet; V, valley glacier; -l, land-terminating; -m, marine-terminating.

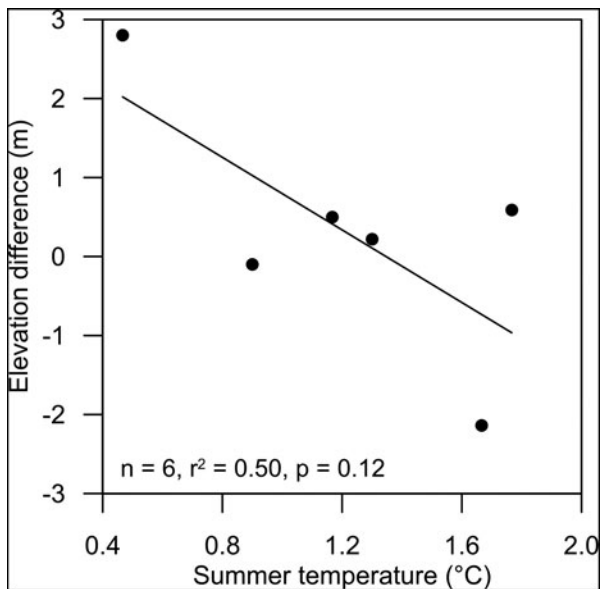


Fig. 8. ICESat tracks/current glacier surface elevation difference vs mean summer (DJF) air temperature on Esperanza meteorological station during the period 2003–2008.

increase in the surface area. The reduced retreat is in line with the recent observations and mass-balance studies of glaciers around the northern AP that show a marked decrease in glacier recession and predominantly positive SMB since the mass-balance year 2009/10.

The ICESat data show that the surface lowering on the IJR-41 ice cap prevailed at least until March 2008. The lowering trend indicated by satellite altimetry over the period 2003–2008 correlates with the mean summer air temperature in the northern AP, implying the significance of surface ablation for annual mass changes of ice caps in this region. The shift

from the lowering trend to an increase in glacier surface elevation probably occurred after the ablation season 2008/09, which ranks among the warmest summers in the north-eastern AP since the mid-20th century. The onset of the positive surface elevation changes in 2009/10 correlates with the period of surface mass gain that was reported from glaciers around the AP in the period 2009–2015. Since 2015, the IJR-41 ice cap has experienced enhanced lowering as indicated by the DGNS data.

The observed aerial and surface elevation changes of the investigated glaciers indicate that the ice caps on LC are out of equilibrium with the current climate. If a continuation of the 1979–2018 surface-lowering trend is assumed, the IJR-41 ice cap will decay over the next decades to centuries except for the highest part of the dome. According to the 2015–2018 lowering scenario, glacier retreat would occur by an order of magnitude rapidly, and the whole glacier will disappear in 66 ± 19 years.

The combination of the historical ICESat data with the present DGNS measurements has proved to be a valuable tool for improving the interpretation of the surface elevation and mass-balance changes on small ice caps in the AP region. Based on these data, the shift from surface mass loss to mass gain prior to the mass-balance year 2009/10 was excluded on the Ulu Peninsula. The satellite altimetry provides a proxy for surface mass changes on the glacier prior to the initiation of the direct field-based measurements in 2009.

The ICESat data records provide a high-accuracy estimate of glacier surface elevation over the first decade of 21st century. A combination of the ICESat dataset and ground-based DGNS measurements excludes terrain information derived from an aerial survey which in polar areas suffers from the lack of an adequate ground control (Fox and Cziferszky, 2008). It thus offers an alternative and independent source of information on glacier thinning in the area of

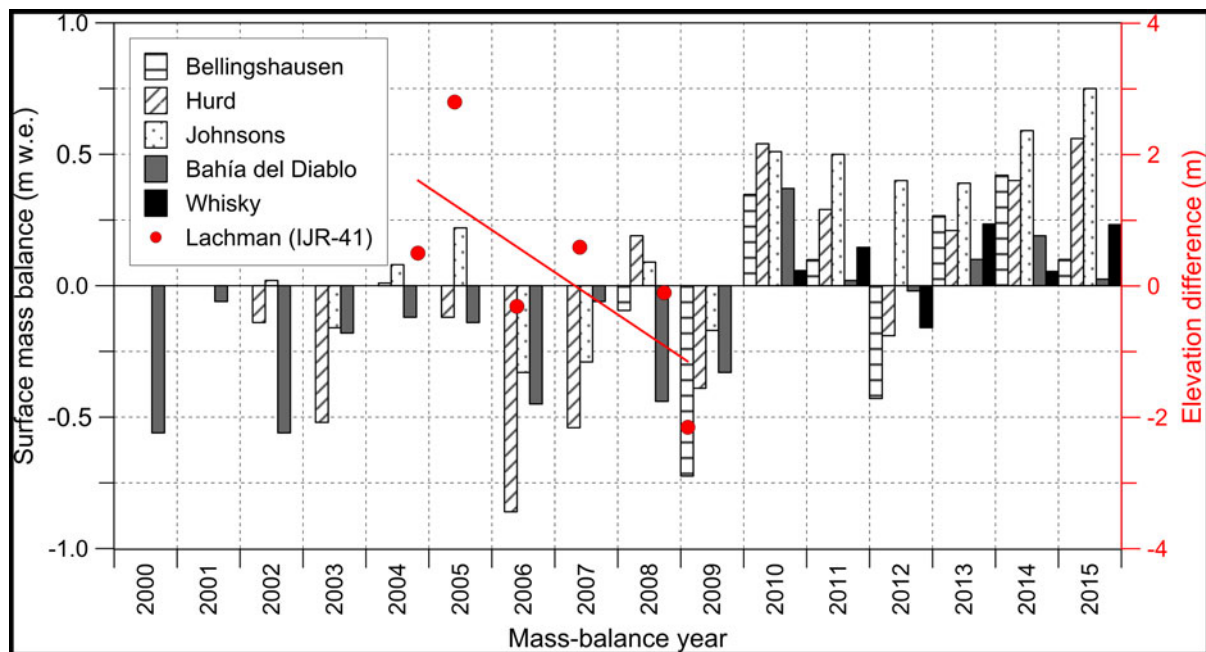


Fig. 9. Surface lowering on IJR-41 ice cap (red dots) during the period 2003–2008 and a change from surface mass loss to mass gain after 2009 as indicated by the ICESat data and mass-balance records for glaciers in the northern part of the Antarctic Peninsula. The right axis indicates the elevation difference between ICESat tracks and current glacier surface. Data adopted from Mavlyudov (2016), Engel and others (2018), WGMS database (2018) and this study.

the AP. The approach applied in our study is limited to glaciers with low-relief topography, especially ice caps and ice domes, where laser altimetry is not biased by steep slopes and surface undulations (Hilbert and Schmillius, 2012). Moreover, it is not realistic to couple the available ICESat measurements over large areas of the AP with DGNSS survey because fieldwork in the hardly accessible areas within this region can be extremely difficult and expensive. As soon as an accurate and homogeneous DEM will be available for the area, it will be possible to evaluate the ICESat measurements in an accurate way.

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